# COMPUTATION OF DEPTH OF PRECIPITABLE WATER IN A COLUMN OF AIR

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It has been found that one of the most significant quantities in hydrometeorological studies is Wp, the depth of precipitable water in a column of air. After some experimentation, the following simple tabular method of computing the value of Wp from aerological observations was developed for use in the Hydrometeorological Research Section.

Let  $M_v$ =total mass of water vapor,

h = height,

 $\rho_w$  = density of water vapor,

 $\rho = \text{density of air},$ 

g=acceleration of gravity,

q = specific humidity, P = pressure in mb,

 $W_p$ =depth of precipitable water (inches).

For a column of water vapor with the horizontal crosssection considered as unity:

$$dM_{v} = \rho_{v} dh \tag{1}$$

$$M_w = \int_0^h \rho_w dh \tag{2}$$

$$dp = -\rho g dh \tag{3}$$

combining (2) and (3)

$$M_{w} = -\int_{P_{0}}^{P_{\lambda}} \frac{\rho_{w} dp}{\rho g} \tag{4}$$

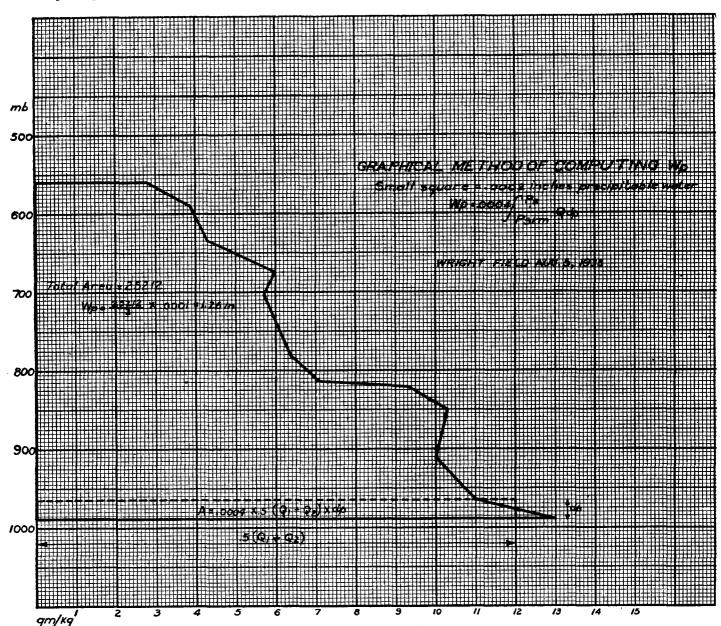


FIGURE 1.—Graphical method of computing precipitable water.

g can be considered a constant with negligible error.

$$M_{w} = -\frac{1}{g} \int_{P_{0}}^{P_{h}} \frac{\rho_{w} dp}{\rho} \tag{5}$$

since  $q = \frac{\rho_w}{\rho}$ 

$$M_{w} = -\frac{1}{g} \int_{P_0}^{P_h} q dp \tag{6}$$

Reversing direction of integrations

$$M_w = \frac{1}{g} \int_{P_0}^{P_0} q dp \tag{7}$$

Pigure Z.

COMPUTATION OF DEPTH OF PRECIPITABLE WATER

(Ohio) Station Wright Field Elevation	Computed by SRS Date December 2. 1938.
Date of Observation August 5, 1935.	Checked by JRR Date December 2, 1938.
Type of Instrument Aerometeorograph	File No. 93508

evels	e,	P	q		1		_	
Sfc.			A	dwp				
_1.		965	11.0	24	24.0	576		
2.		911	10.0	54	21.0	1134		
3.		851_	10.3	60	20.3	1218		
4.		824	9.3	27	19.6	529		
5.		814	7.1	10	16.4	164		
6.		776	6.3	38_	13.4	509		
7.		705	5.7	71	12.0	852		
8.		675	6.0	30_	11.7	351		
9.		633	4.5	42	10.3	435		
10.		692	3.9	41	8.2	536		
11.		562	2.8	30	6.7	201		
12.								
13.		\ <del> </del>						
14.			LI					
15.		<u> </u>						
16.					L			
17.		<u></u>						
18.								
19.								
20.								
			JL		Total	6303	1.26	

e - wapor pressure in mb.

dp = P1 minus P2, etc.

P = pressure in mb.

On was a day etc.

q - specific humidity g/kg.

A - Qdp

0.0002 A = Wp, depth in inches of precipitable water.

Dimensional formulas may be expressed as follows:

q=non-dimensional ration.

$$P=\frac{ML}{L^2T^2}$$
.

$$g=\frac{L}{T^2}$$

$$\frac{Pq}{g} = \frac{MLT^2}{L^2T^2} = \frac{M}{L^2}.$$

In the C. G. S. system this could be expressed as  $\frac{gm}{cm^2}$ .

Since 1 cc. of water has a mass of 1 gram this becomes  $\frac{\text{cm}^3}{\text{cm}^2}$ 

or cm. In order to evaluate the constant which would make the result expressible in inches of precipitable water, the following substitutions must be made:

g=980 cm/sec<sup>2</sup>.

1 mb=1,000 
$$\frac{\text{dynes}}{\text{cm}^2}$$

1 inch=2.54 cm.

1 gm./kg. = 
$$\frac{1}{1000}$$
 gms./gm.

$$\frac{1000}{980 \times 2.54 \times 1000} = .0004.$$

Figure 3 COMPUTATION OF DEPTH OF PRECIPITABLE WATER

Station Nurfreedore Stevation 174 M. Computed by SRS Date December 5, 1936.

Date of Observation Mrv 25, 1935 Checked by JRR Date December 5, 1936.

Typo of Instrument File No.

	Lovels	9.0	P 1000	9 5.6	dp	Q2m	<del>-</del>	
			1			- <del>  42m</del>   -		dwp
500 E.	1.	7.9	964	5,1	36	10,7		.077
1000 N°	2.	8,4	907	5.5	57	9.5		.108
500 ¥.	3.	5.2	850	3.8_	57_	8.2		093
.¥ 000	4.	4.0	799	3.1	_ 51	6.9		.084
500 N.	5.	2.4	752	2.0	_47_	5.2		.048
000 ¥.	6.	1.9	706	1.7	46	3.7		.034
000 N.	7.	0.9	621	0.9	86	2.6		.044
000 N.	8.	_0.6	547	0.7	74	1.6		.024
	9.		<u> </u>			1		
	10.							
	11.					1		
	12.							<del>                                     </del>
	13.							
	14.					T		
	15.				_			<del>                                     </del>
	16.							
	17.							
	18.							
	19.							
	20.							1
						Total		.51

e - vapor pressure in mb.

dp = F1 minus P2, etc.

P = pressure in mb.

Q<sub>2m</sub> = q<sub>1</sub> + q<sub>2</sub>, etc.

q = specific humidity g/kg.

A = Qdp

Therefore  $Wp = .0004 \ q \ dp$  inches. where q is expressed in gm./kg.

and dp in millibars.

Furthermore, it has been found that for practical purposes the precipitable water present from the surface to 5 kilometers above sea level may be considered as equal to the total amount of precipitable water, since the amount above that elevation is very small.

0.0002 A . Wp. depth in inches of precipitable water.

Thus 
$$Wp = .0004 \int_{P_{\text{Skm}}}^{P_{\text{c}}} q \, dp \tag{8}$$

It is possible to approximate the value of the above integral graphically by plotting significant values for q against P on linear cross section paper and computing the area under the curve.

To illustrate, the following data for Wright Field, Dayton, Ohio, August 5, 1935, were thus plotted (fig. 1).

q			
13. 0 11. 0			
10. 0 10. 3			
9.3 7.1			
6. 3 5. 7			
6. 0 4. 3			
3. 9 2. 8			

 $P_n$ ,  $P_{n-1}$ =pressure at any significant level and pressure at next lower significant level, respectively.

 $Q_n$ ,  $Q_{n-1}$  = specific humidity at  $P_n$  and  $P_{n-1}$ , respectively.

From figure 1 it will readily be seen that the value of  $A_n$ ,  $_{n-1}$ , is equivalent to the area of a rectangle whose sides are

$$P_{n-1}-P_n \text{ and } .5(Q_{n-1}+Q_n)$$

$$A_n, \ _{n-1}=.5(P_{n-1}-P_n)(Q_{n-1}+Q_n)$$
and  $Wp_n, \ _{n-1}=.5\times.0004(P_{n-1}-P_n)(Q_{n-1}+Q_n)$ 

$$=.0002(P_{n-1}-P_n)(Q_{n-1}+Q_n)$$
and  $Wp = \sum .0002 (P_{n-1}-P_n)(Q_{n-1}+Q_n)$ 

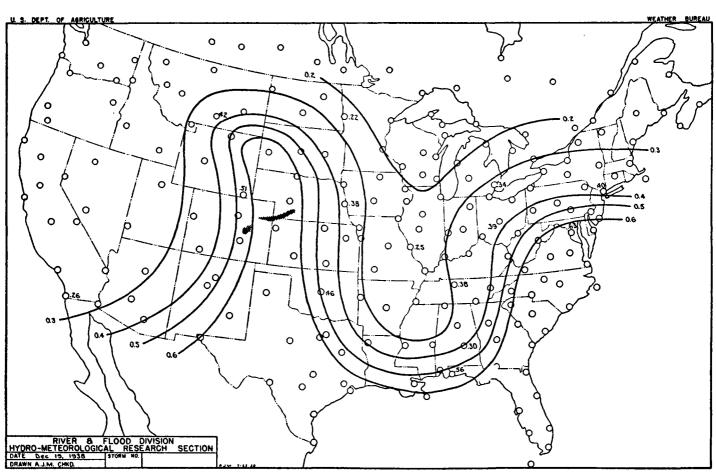


Figure 4.—Distribution of precipitable water from 2 to 5 kilometers on the morning of May 31, 1935, during the intense rain which caused the Republican River flood. Shaded portion shows the area of flood rains.

The scale in figure 1 was chosen for convenience. Since one small scalar element equals 0.1 gm./kg. and 5 millibars, it will be seen that one small square of area is equivalent to  $\frac{1}{2}$  equation units. Thus if  $A_t$  represents the total number of small squares, then

$$\frac{.5\times4}{10,000}A_i$$
=inches precipitable water,

or 
$$\frac{.0004}{2}A_i$$
=inches precipitable water.

Let  $A_{n,n-1}$ =an area on the graph between any two significant pressures expressed in equation units.  $(p=1 \ mb., q=1 \ gm./kg.)$ .

If, instead of being plotted, the data were entered on a suitable form, the result is obtainable by means of very simple calculations. Figure 2 shows such a form.

In column P are entered pressures at significant levels. In column Q, specific humidities for corresponding levels. In column dp are the successive values for  $P_{n-1}-P_n$ . In column  $Q_{2m}$  are the successive values for  $Q_{n-1}+Q_n$ . Column A is the product of dp and  $Q_{2m}$ . The final result is the sum of column A multiplied by the constant .0002. In some great the values of  $W_n$  between periods standards

In some cases the values of Wp between various standard elevations are required. For this purpose it is more convenient to use standard level data from Weather Bureau Form 1109, instead of significant level data. The error involved in using standard level data is very small.

For this purpose column dWp has been provided. This column is simply the product of A and the constant .0002. In actual practice it is unnecessary to use both column A and dWp together, since dWp may be entered directly by multiplying dp by twice column  $Q_{2m}$  and setting decimal places properly. Three decimal places are sufficient. Column dWp may then be summed up and the third decimal place dropped. This will give Wp directly (fig. 3).

Column e is provided for cases where specific humidity is not given, but must be computed from vapor pressure, e.

It is not the purpose of this article to discuss the use of the quantity Wp. However, a chart showing the daily distribution of precipitable water from 2 to 5 kilometers over the United States is herein presented for general interest.

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